



## Uncertainties in models for glacial isostatic adjustment

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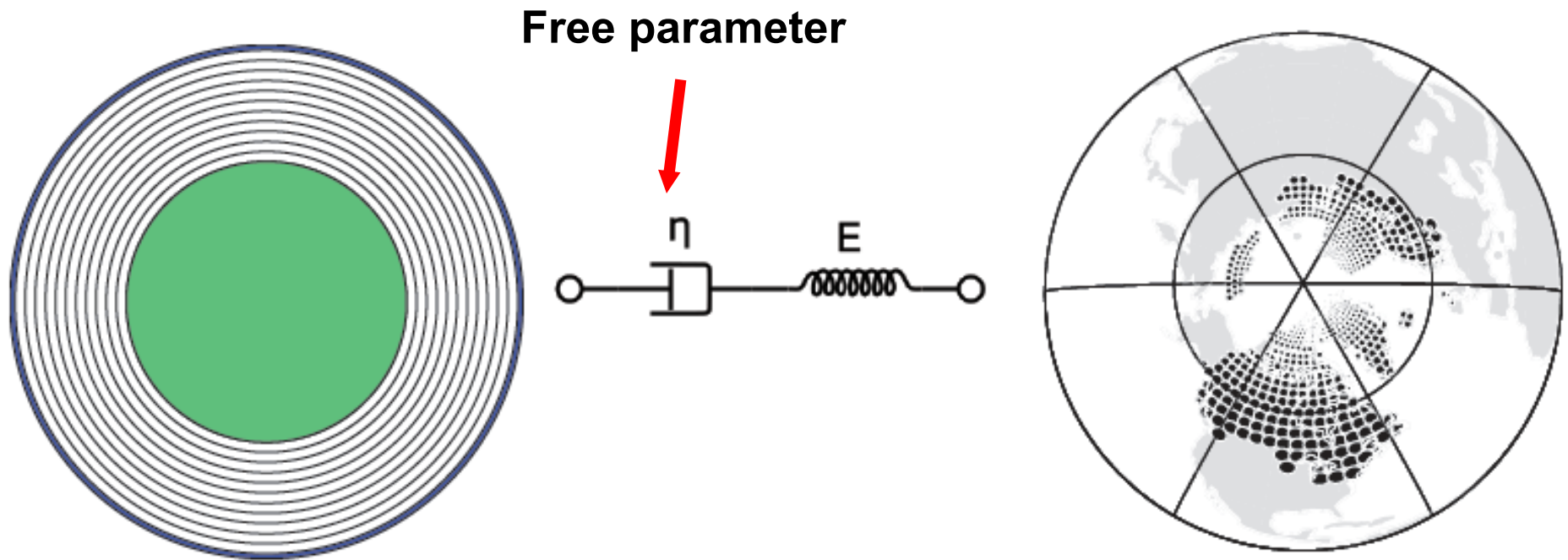
# Uncertainties in models for glacial isostatic adjustment

Wouter van der Wal & Valentina Barletta  
GGFC Workshop  
Vienna, April 20, 2012

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# Standard GIA model

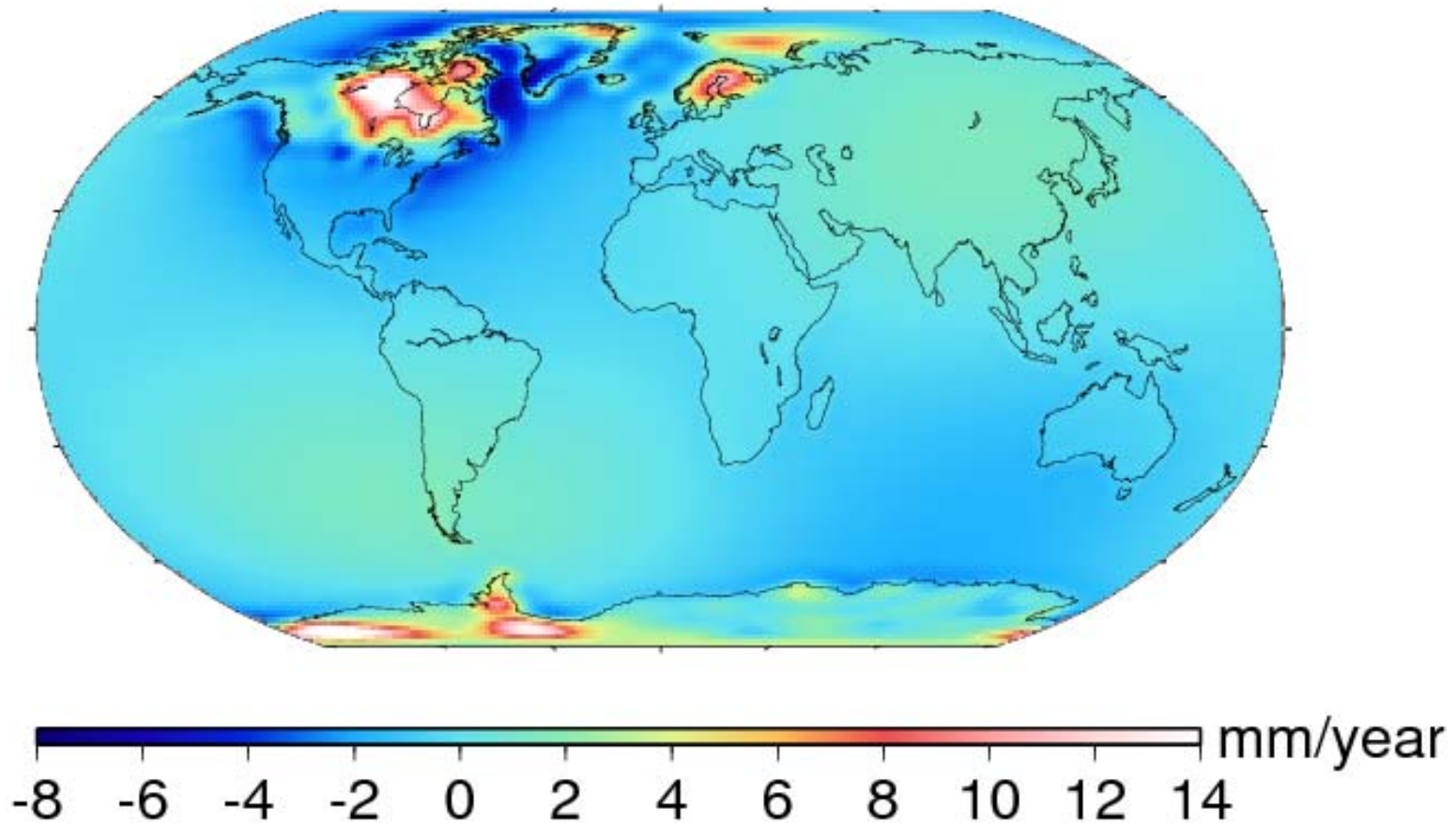


From: Paolo Stocchi, IMAU Utrecht

# STANDARD MODEL

# Standard GIA model

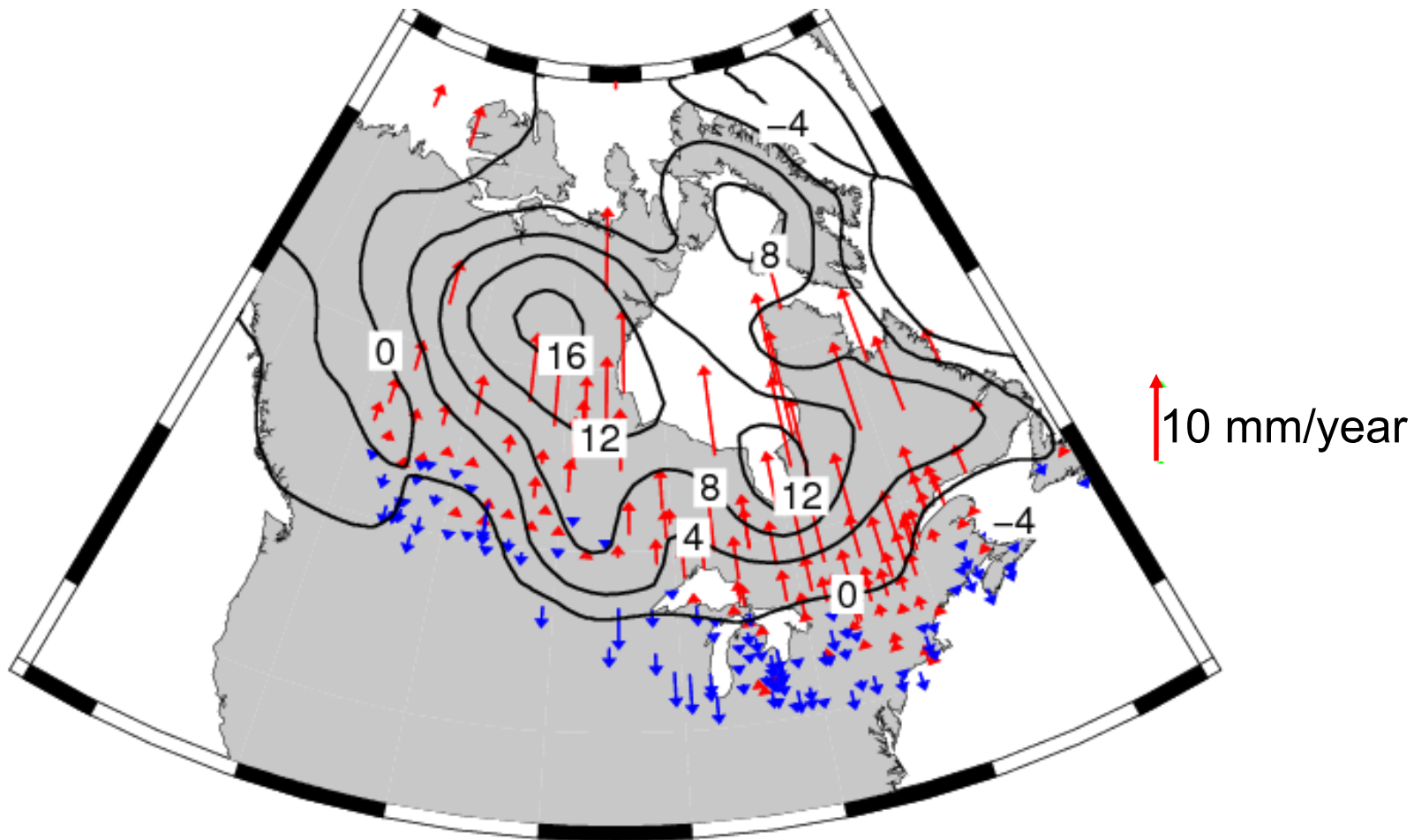
Uplift rate ICE-5Gv1.2/VM2



Uplift rate from Peltier submission to Special Bureau for Loading website

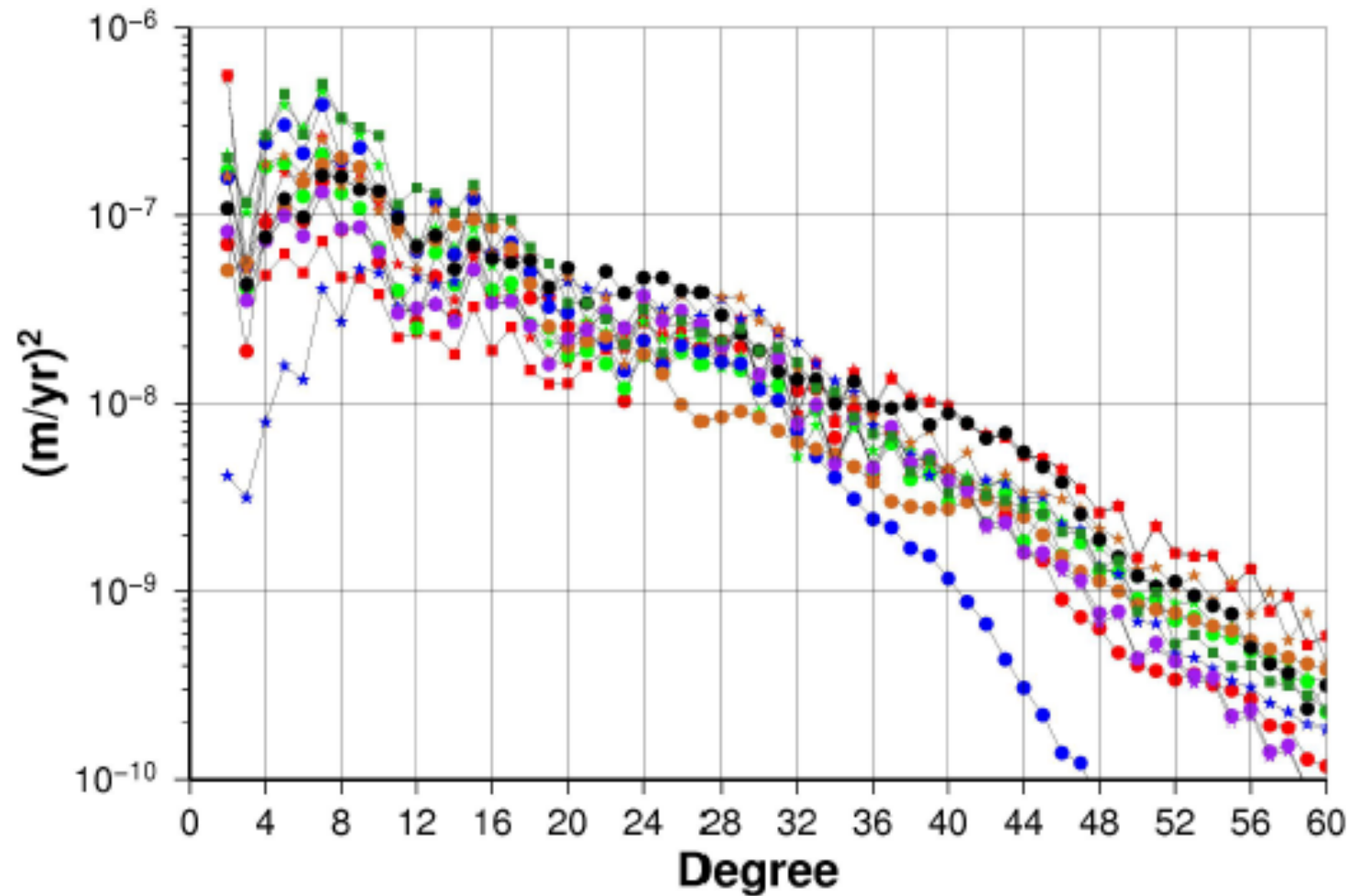
# Standard GIA model

**Contours:** ICE-5G/VM2 **Arrows:** GPS uplift rates Sella et al. (2007)



van der Wal et al. (Canadian Journal of Earth Sciences 2009)

## Uplift Rate Spectrum



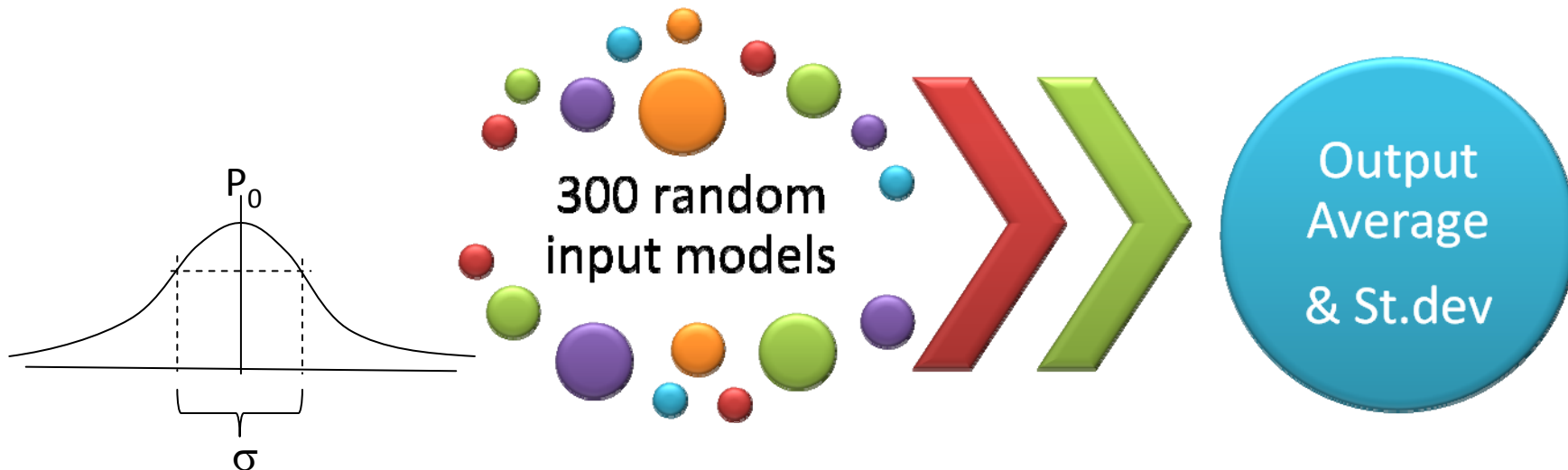
Guo et al, J. Geodyn. (2012)



# Method for uncertainty propagation

## Monte Carlo Method:

hundreds of randomly generated input models with a Gaussian distribution with selected sigma around the input reference model  $P_0$

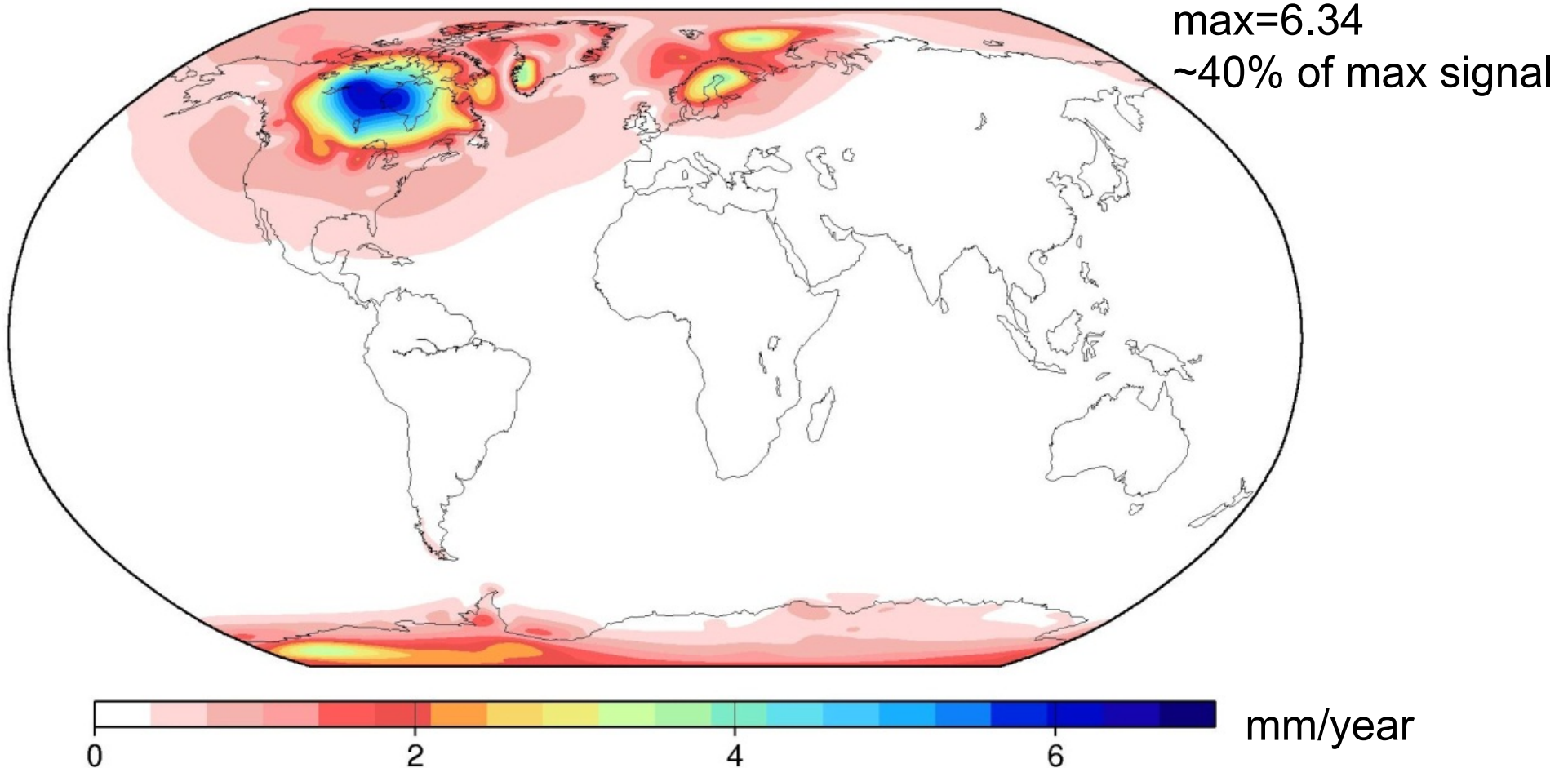


## Reference model:

Ice and Earth model: ICE5G (incompressible - 5Layer - VM2 – L90)

# Uncertainty propagation: viscosity

Standard deviation



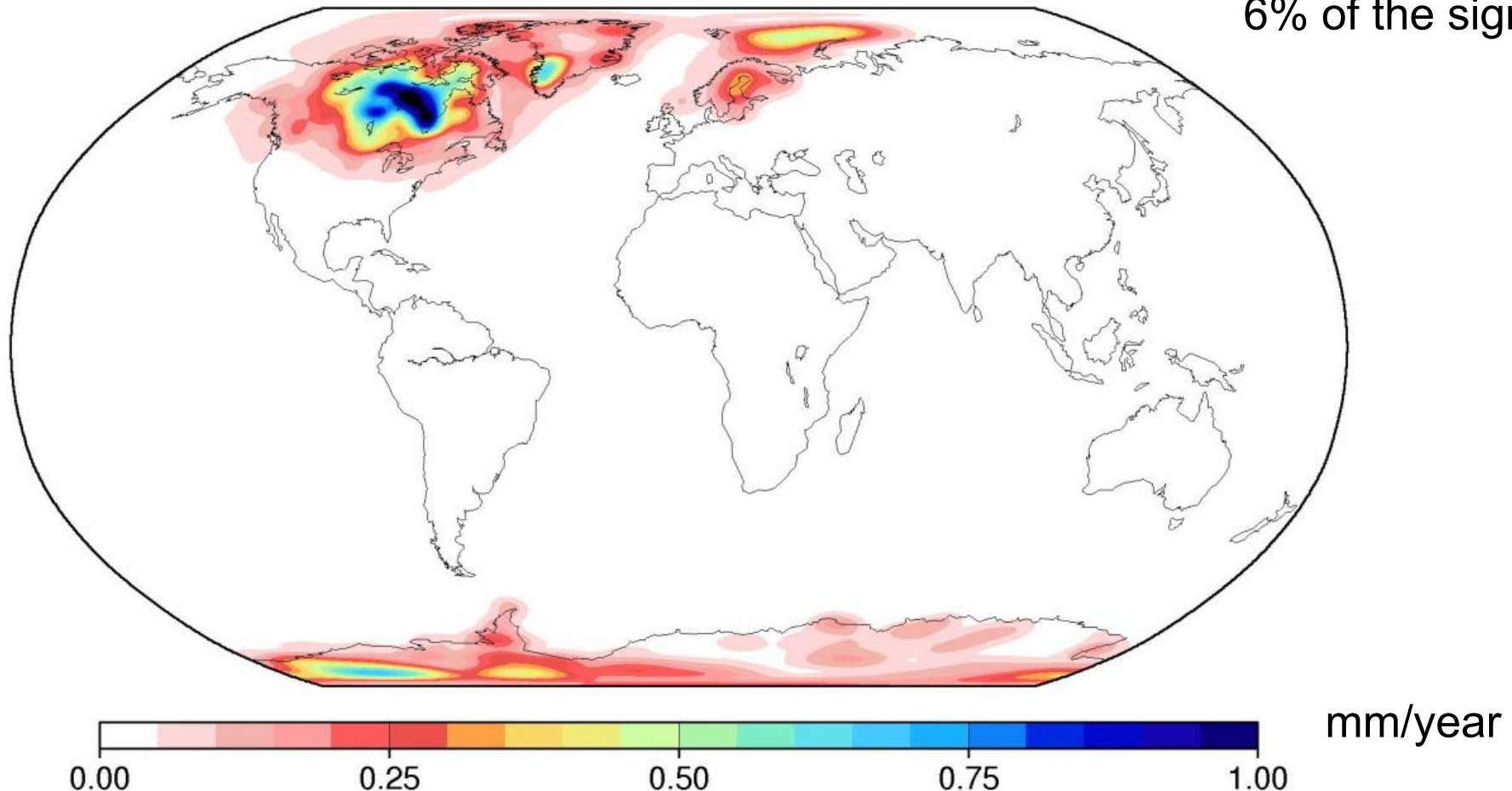
Variation of the viscosity by  $\pm 0.3$  in Log10 scale, i.e. by  $10^{\pm 0.3}$

# Uncertainty propagation: ice height

Standard deviation

Max = 1.1

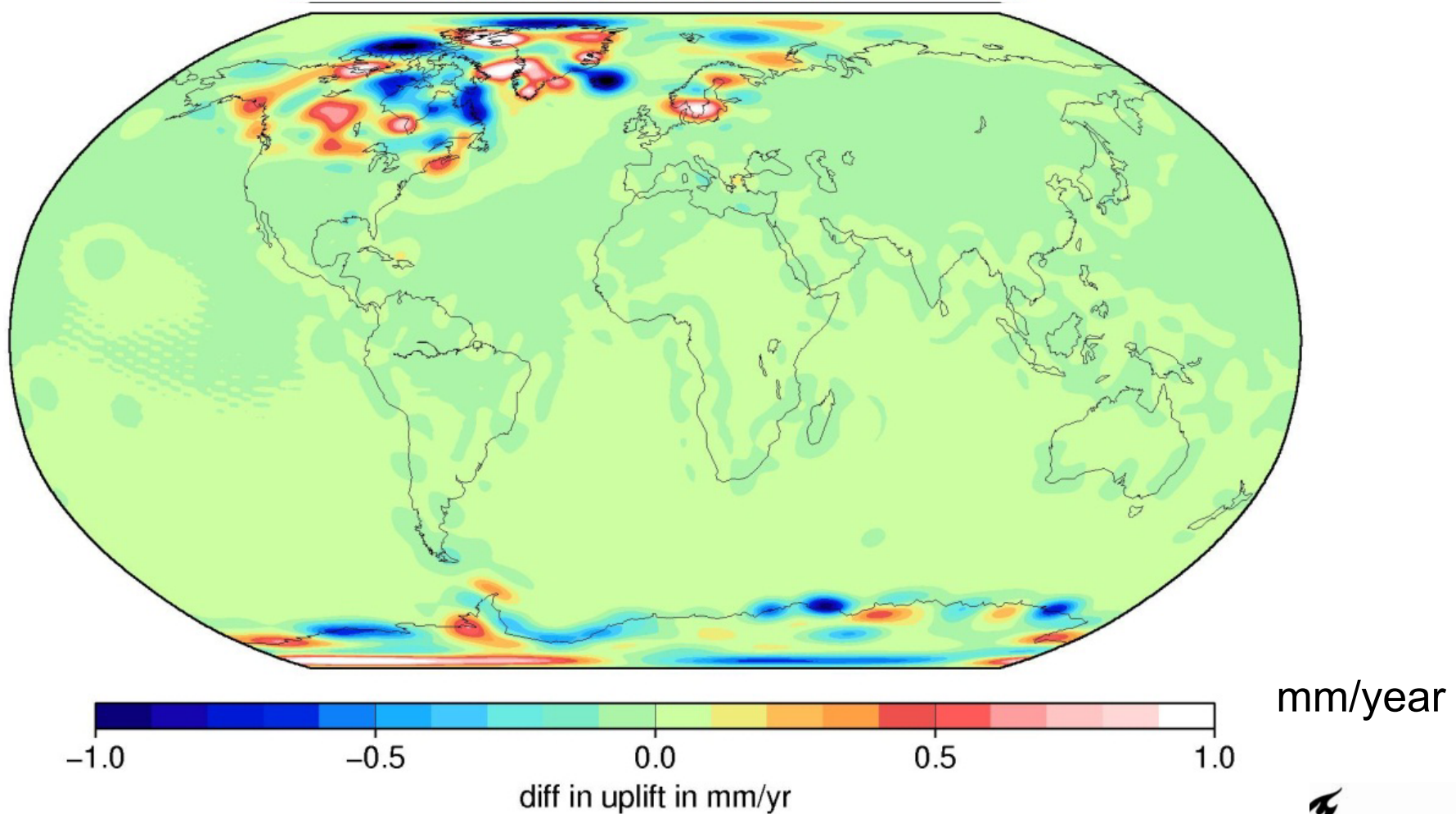
6% of the signal



**I10:** Variation of  $\pm 30\%$  of the Ice thickness for each time and location. Where  $I(t, w)$  is the same as today, we assumed a  $\pm 10\%$  variation for the ice  $< 800\text{m}$ .

# Uncertainty: implementation

Difference in uplift rate

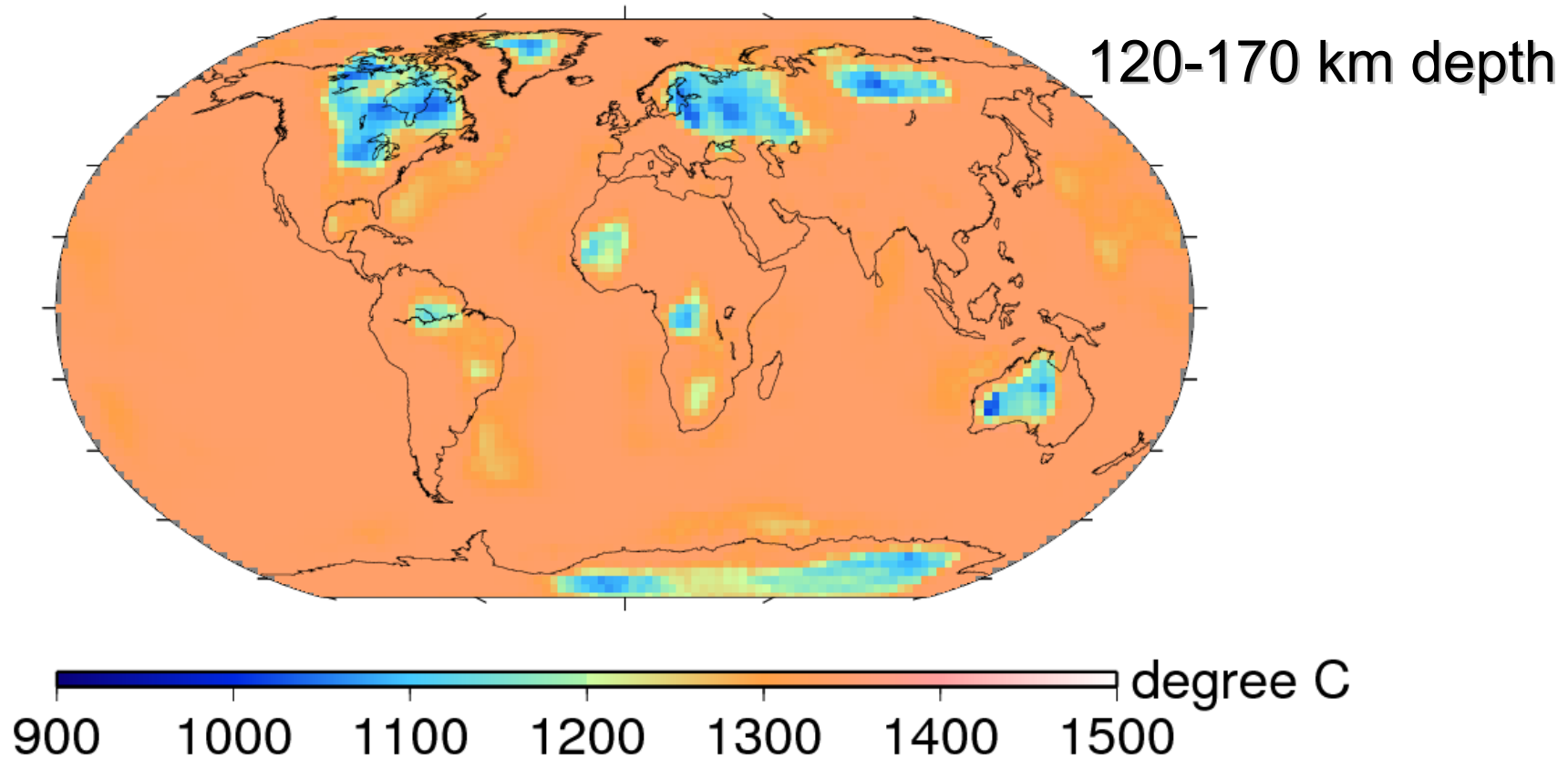


Same Ice model ICE5G, Same Earth model VM2

Difference in initial sampling of the ice model and the ocean function

# FINITE-ELEMENT MODEL

# Uncertainty: lateral variation



- Heatflow measurements extrapolated by a global seismic model  
(Shapiro & Ritzwoller 2004)+ heat diffusion equation

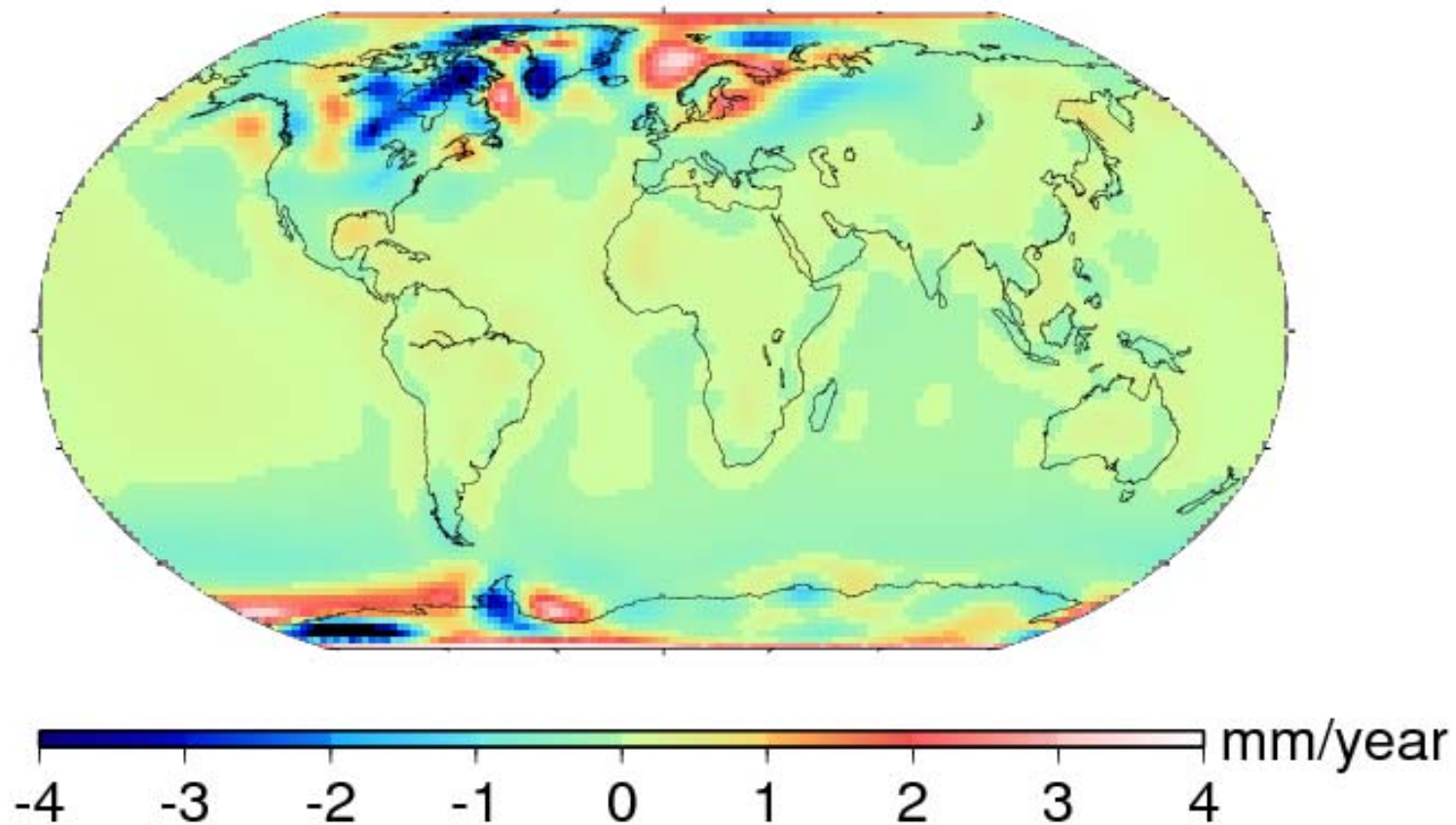
van der Wal et al., (in prep.)



# Uncertainty: lateral variation

Lateral varying – ICE-5G/VM2

-6.8 mm/year



van der Wal et al., (in prep.)

# Uncertainty: mantle deformation

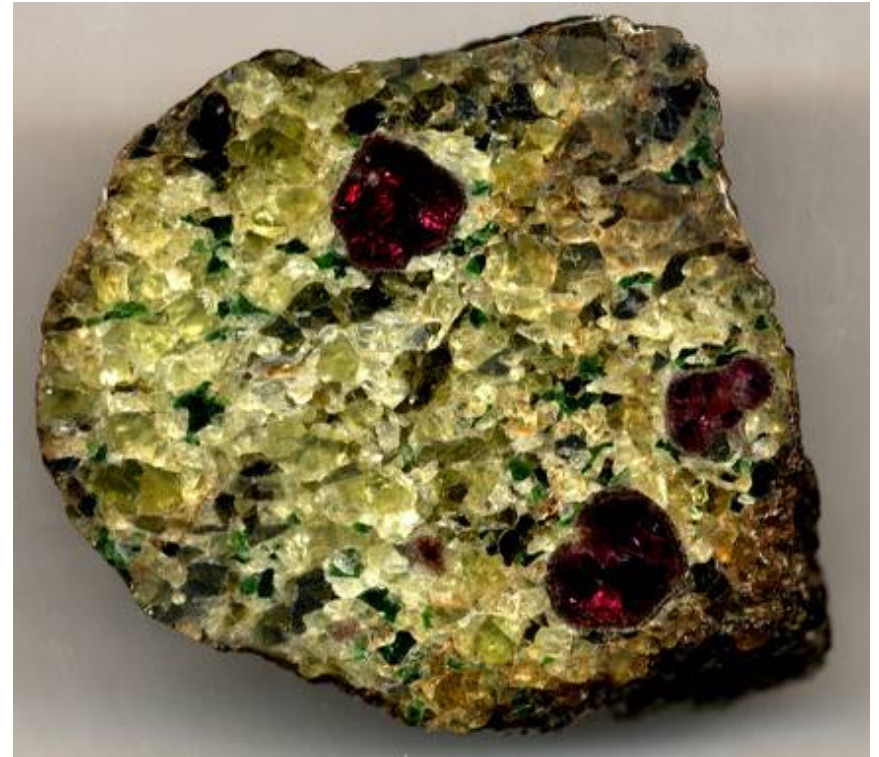
Mantle rocks in the laboratory

$$\dot{\tilde{\epsilon}}_{ij} = \left( \frac{3A_{n=1}}{2} + \frac{3}{2} A \tilde{q}^{n-1} \right) S_{ij}$$

$S_{ij}$  *deviatoric stress tensor*

$\tilde{q}$  *von Mises equivalent stress*

$n$  *stress exponent (3.5)*



From: Martyn Drury, Utrecht Univ.

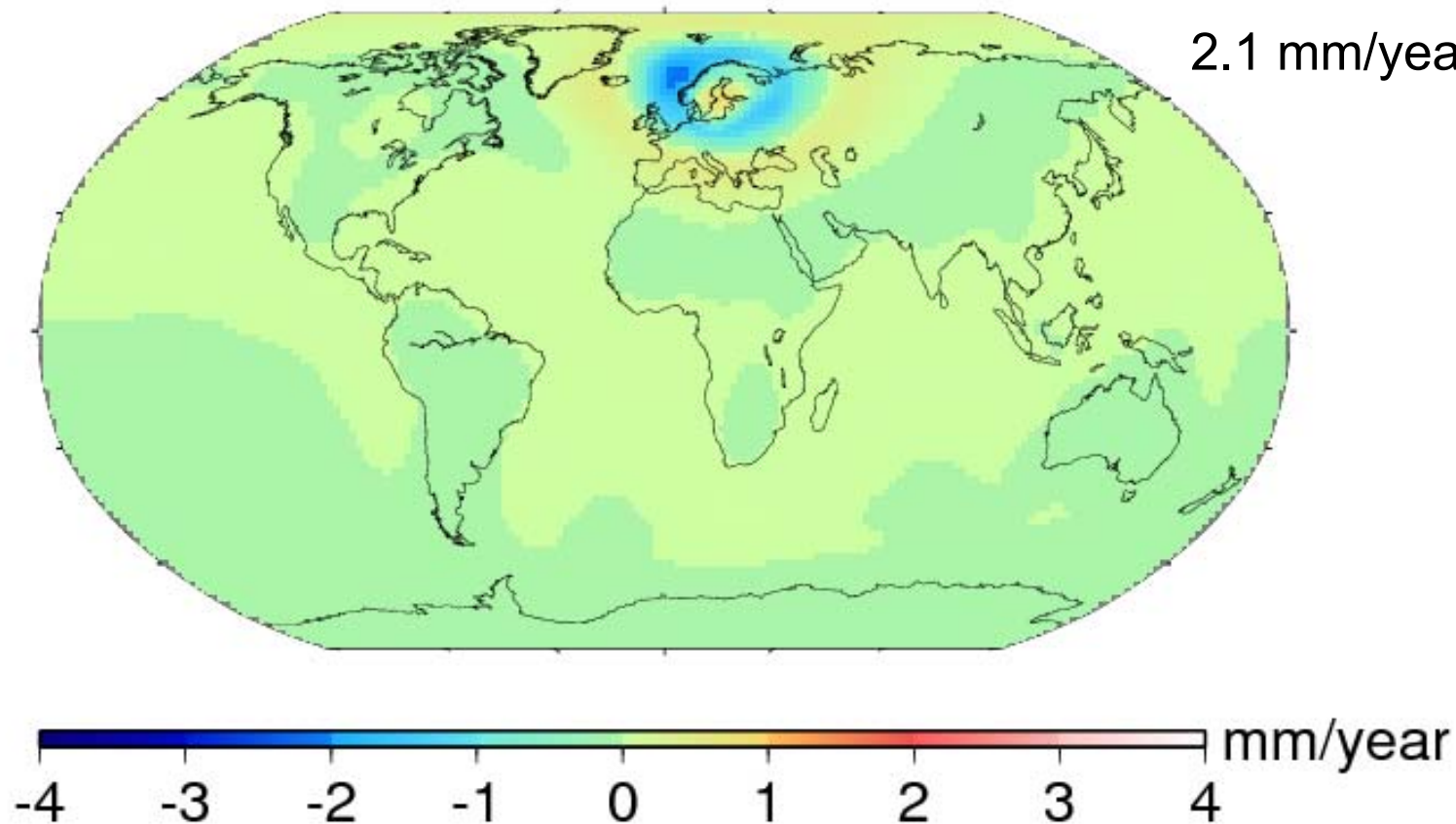


# Uncertainty: mantle deformation

Stress-dependent flow law

Max.

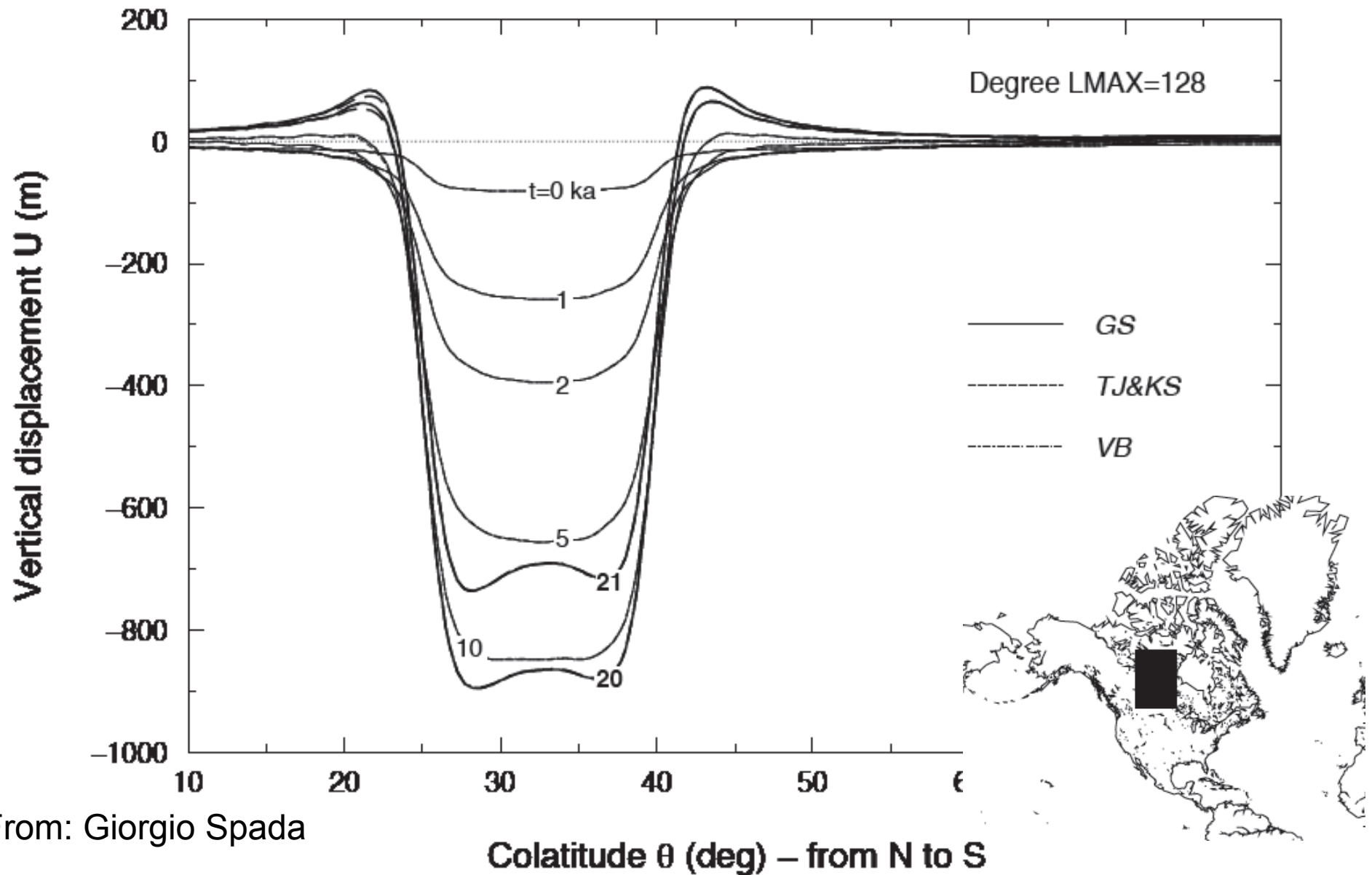
2.1 mm/year



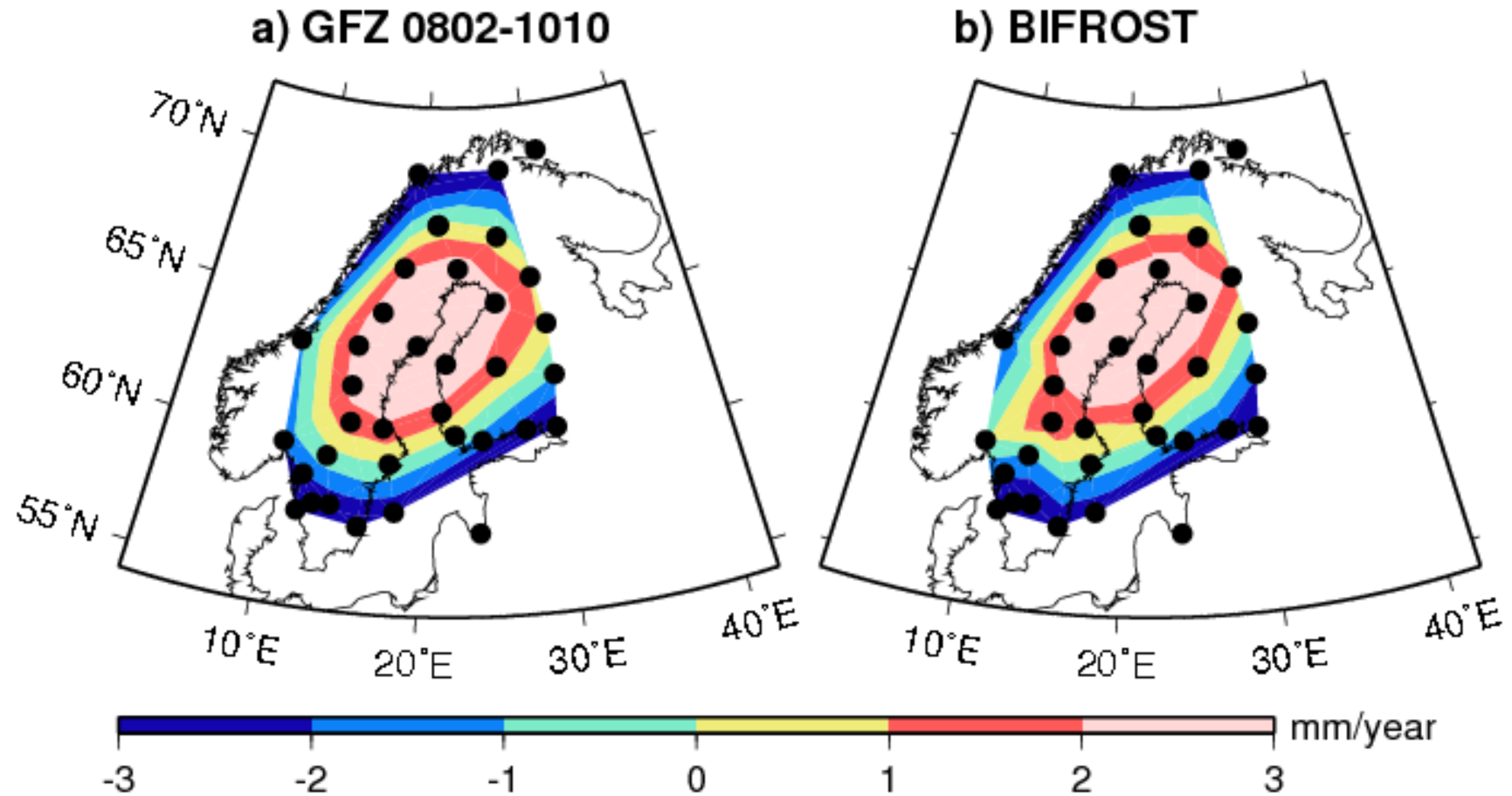
van der Wal et al., (in prep.)

# SOLUTIONS?

# Solutions: Benchmark



# Solutions: Data



Van der Wal et al. (GJI 2011)

# Summary

**Standard model:** Viscosity - 6.3 mm/a, other Earth model – 1.9 mm/a, ice height - 1.1 mm/a, rotational feedback ??

**3D:** 6.8 mm/a, **Flow law:** 2.1 mm/a

## **Solutions:**

- Use uncertainty estimate
- Benchmark
- Use other measurements
- Constrain the model for the region of interest
- Constrain the model with information from other Earth sciences

Funded by TOPO-EUROPE, a EUROCORES project from  
the European Science Foundation

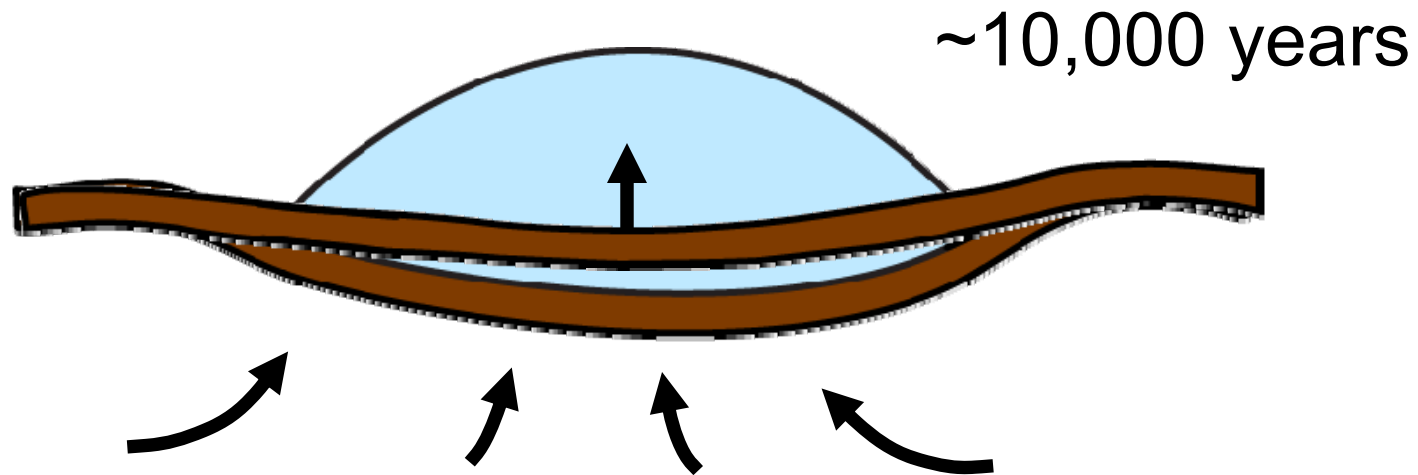
With support by COST Action ES0701 "Improved constraints  
on models of Glacial Isostatic Adjustment"



# BACKUP SLIDES

# Glacial Isostatic Adjustment (GIA)

rising



Flow in the mantle determined by viscosity



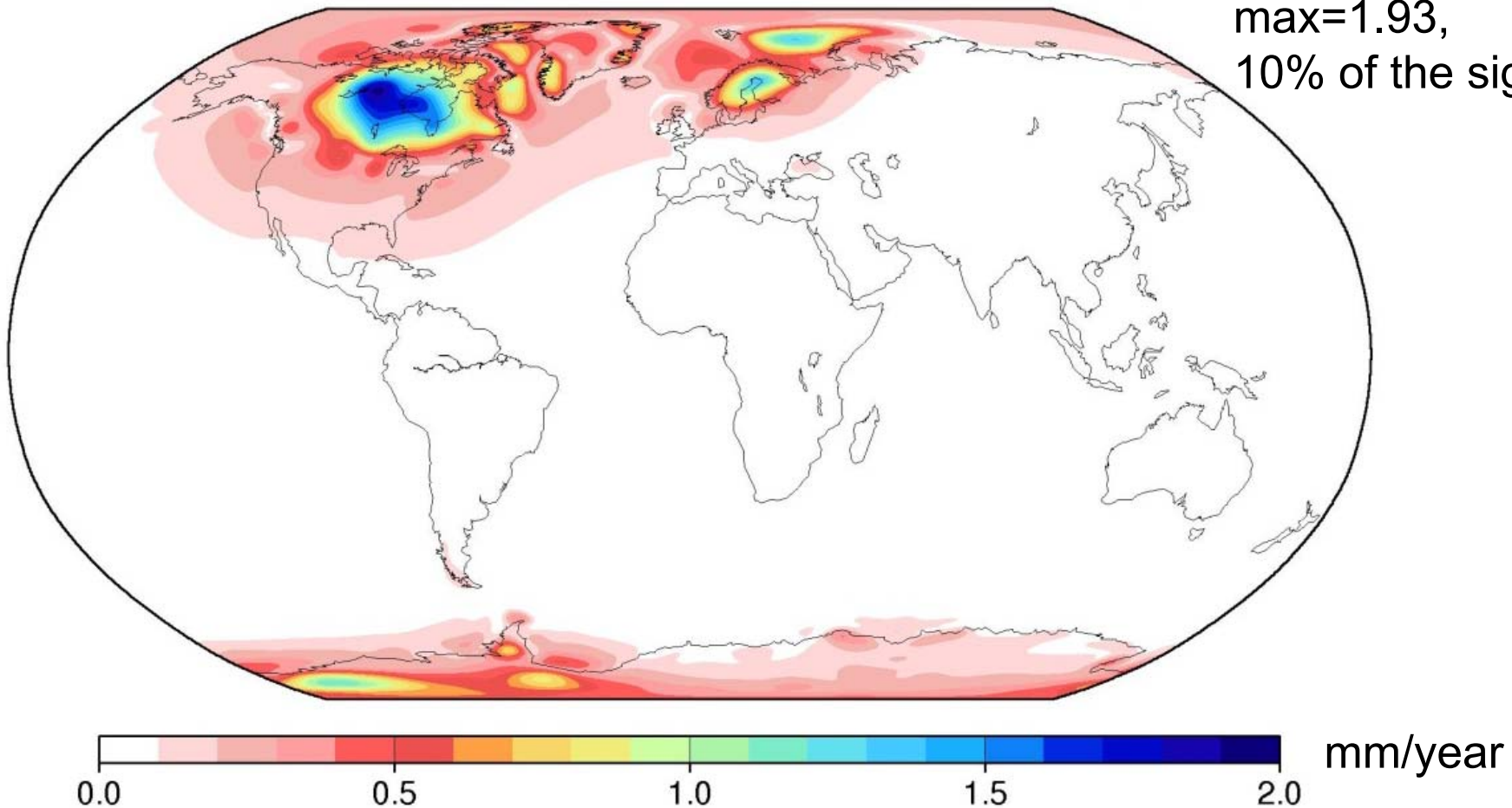
# Results: best fitting mantle viscosities

	$\eta_{UM}$ [ $10^{20}$ Pas]	$\eta_{LM}$ [ $10^{20}$ Pas]
Tushingham & Peltier (1991)	10	20
Mitrovica & Forte (2002)	4	80
Kaufmann & Lambeck (2002)	7	200
Wolf et al. (2006)	3.2	160
Paulson et al. (2007)	5.3	23
GPS (ICE-4G)	8	32
GRACE (ICE-4G)	64	256
Historic sea level (ICE-4G)	16	32
Historic sea level (ICE-5G)	16	256

# Uncertainty propagation: Earth model

Standard deviation

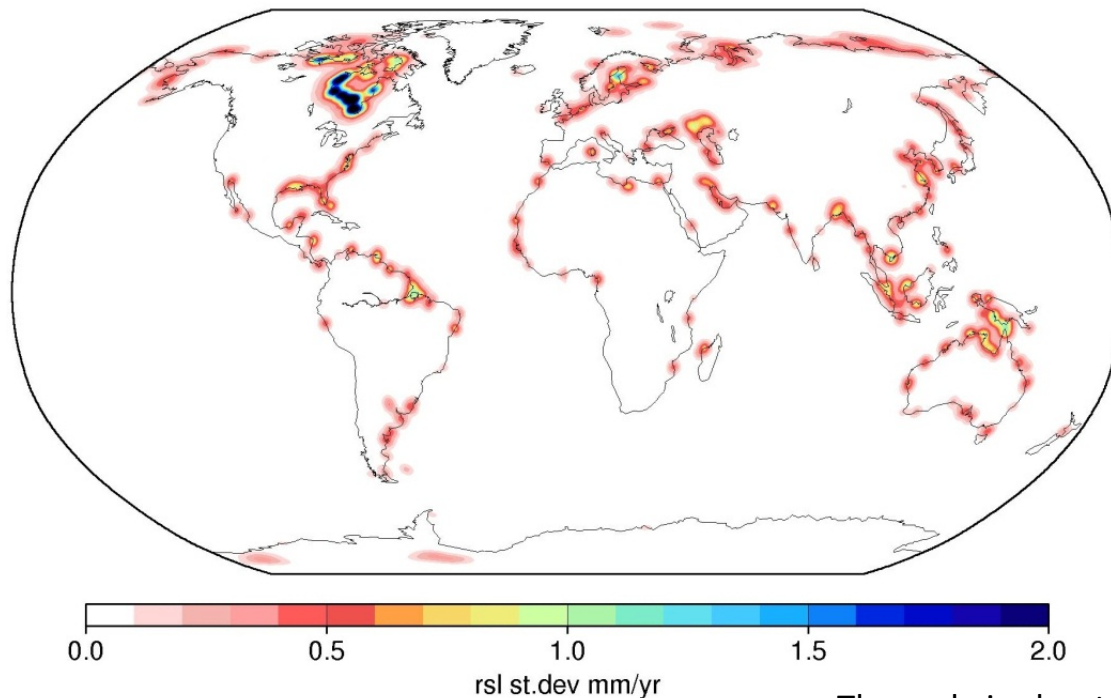
max=1.93,  
10% of the signal



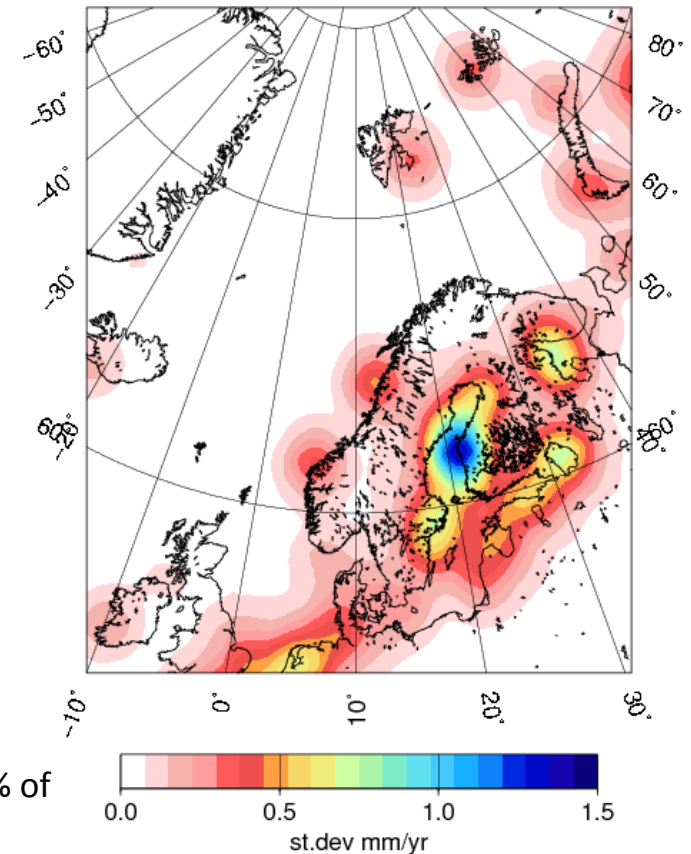
Variation of Lithospheric thickness  $\pm 5$  km, Density  $\pm 5\%$ ,  $V_s \pm 5\%$ , and viscosity by  $\pm 0.05$  in Log10 scale

# Propagation of Ocean Function uncertainties

St. Dev for RSL (GPS and tide-gauges)

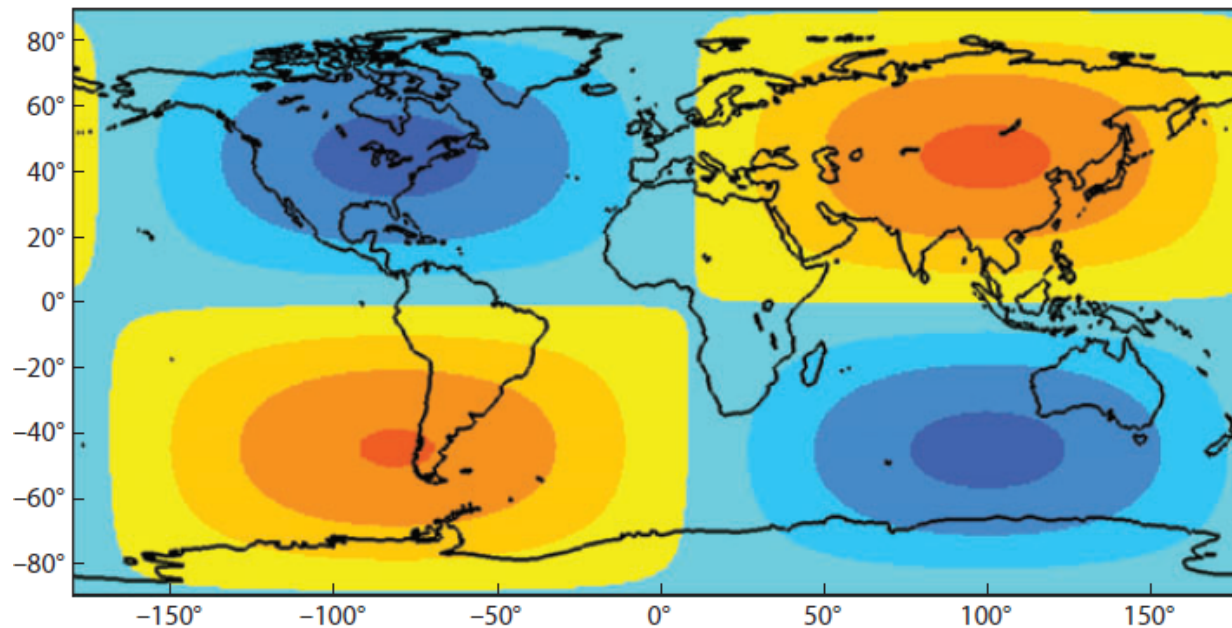


The scale is about 15% of the whole signal.



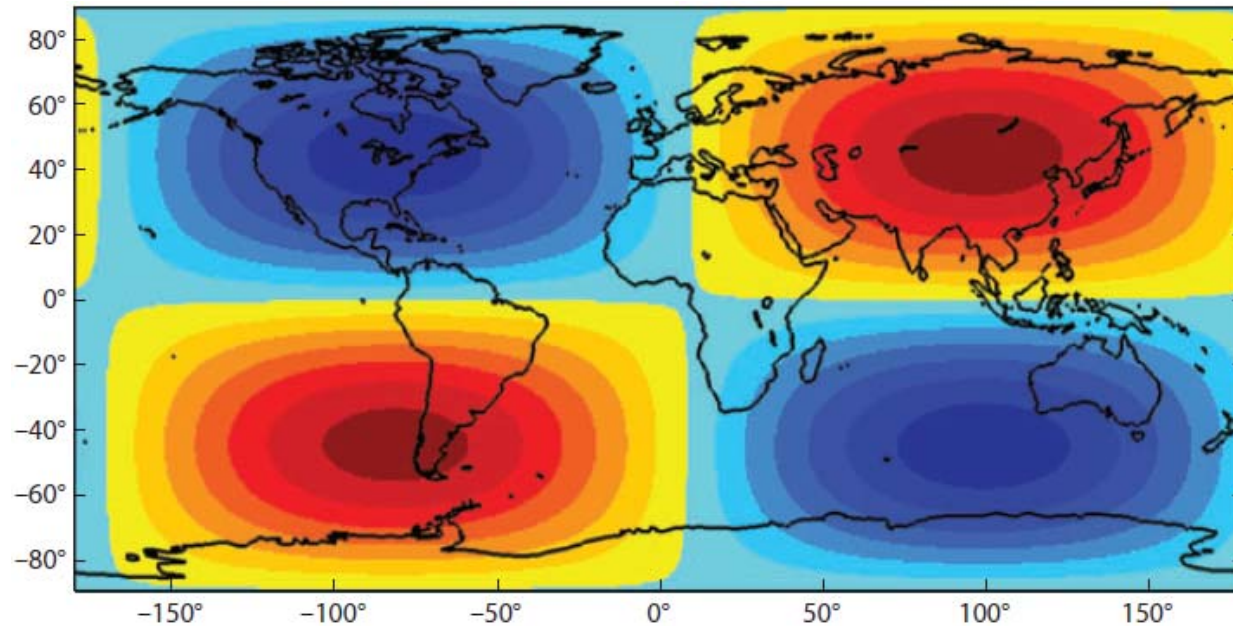
**O2:** Variation of  $\pm 10\%$  of the paleotopography  $T(t, \omega)$  for each time  $t$  and only in locations  $(\omega)$  within a belt following the shorelines. From the paleotopography then we compute the ocean function  $FO(t)$  by setting  $FO = 1$  where the paleotopography is negative, and  $FO = 0$  otherwise.

# Uncertainty propagation: rotational feedback



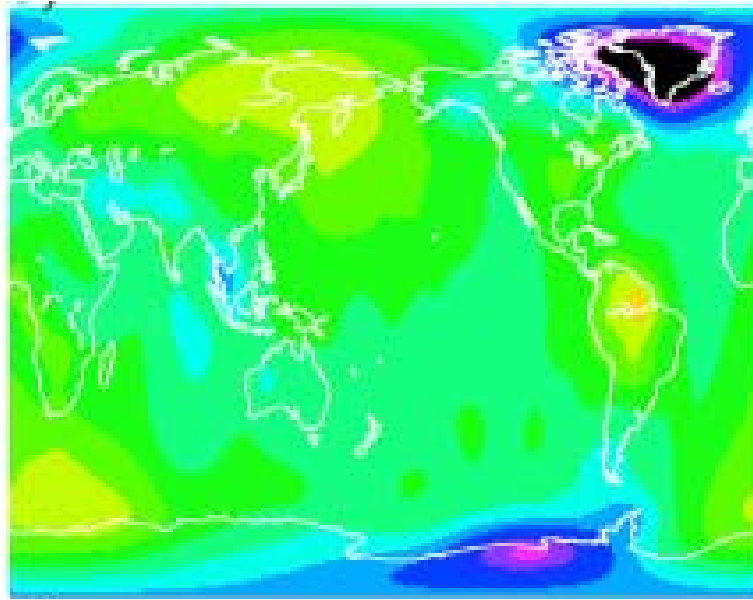
Mitrovica & Wahr, Ann. Rev. Earth Planet. Sci. (2011)

# Uncertainty propagation: rotational feedback

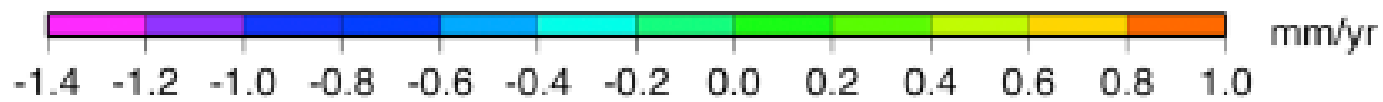
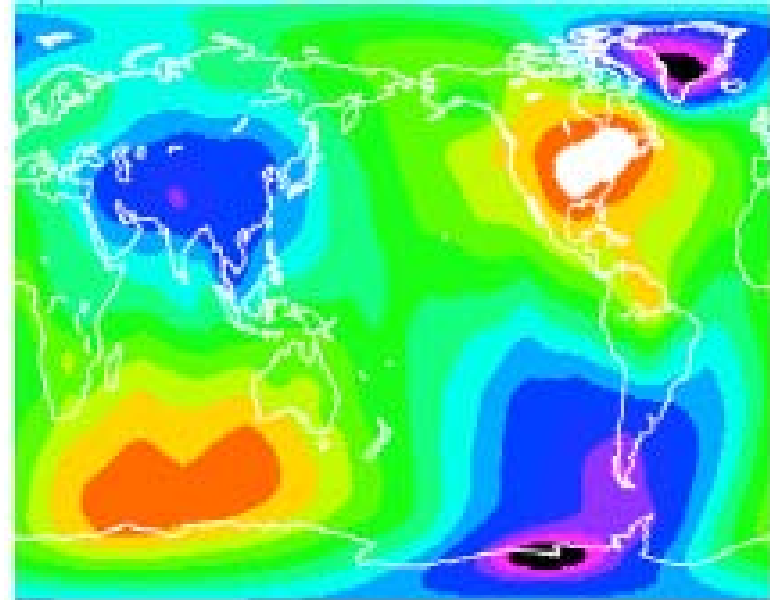


# Uncertainty: rotational feedback

GRACE – Paulson et al (2007)

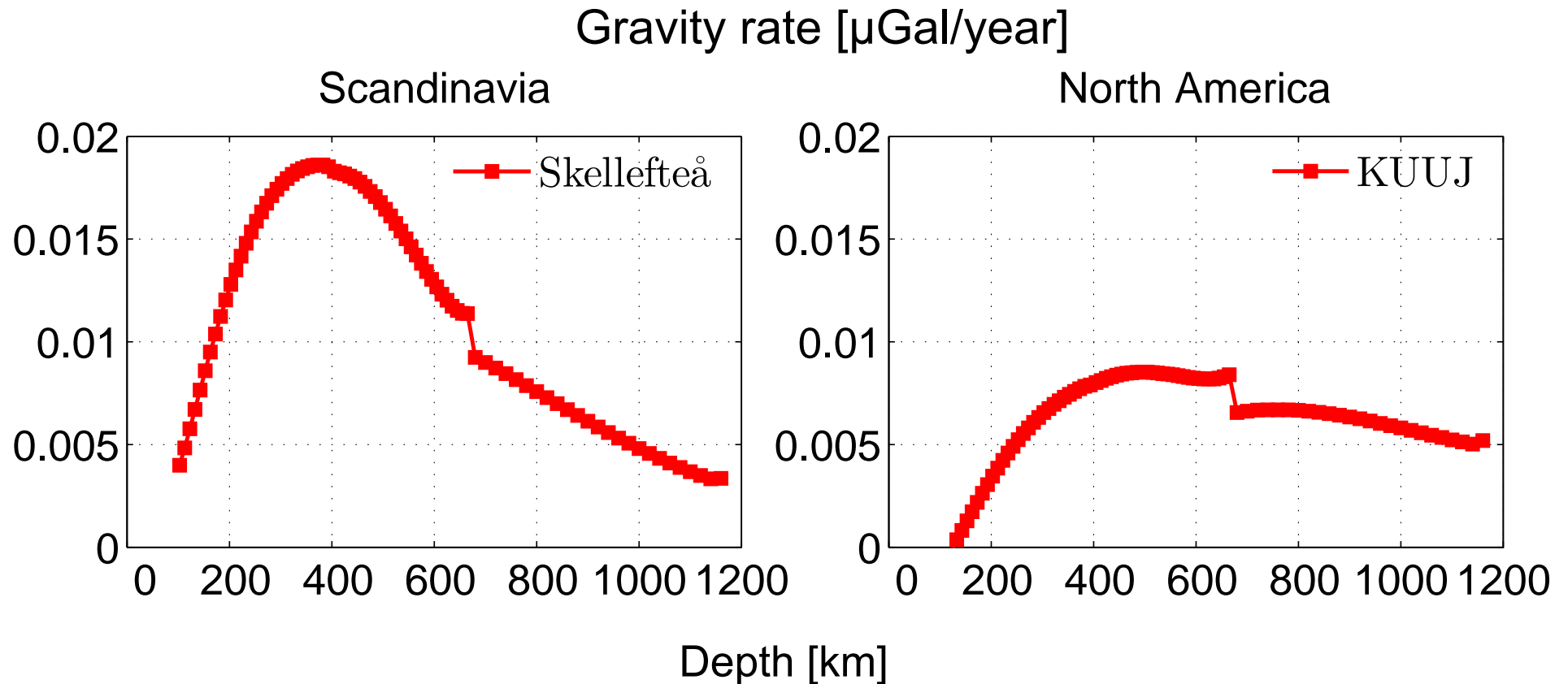


GRACE – Peltier (2004)



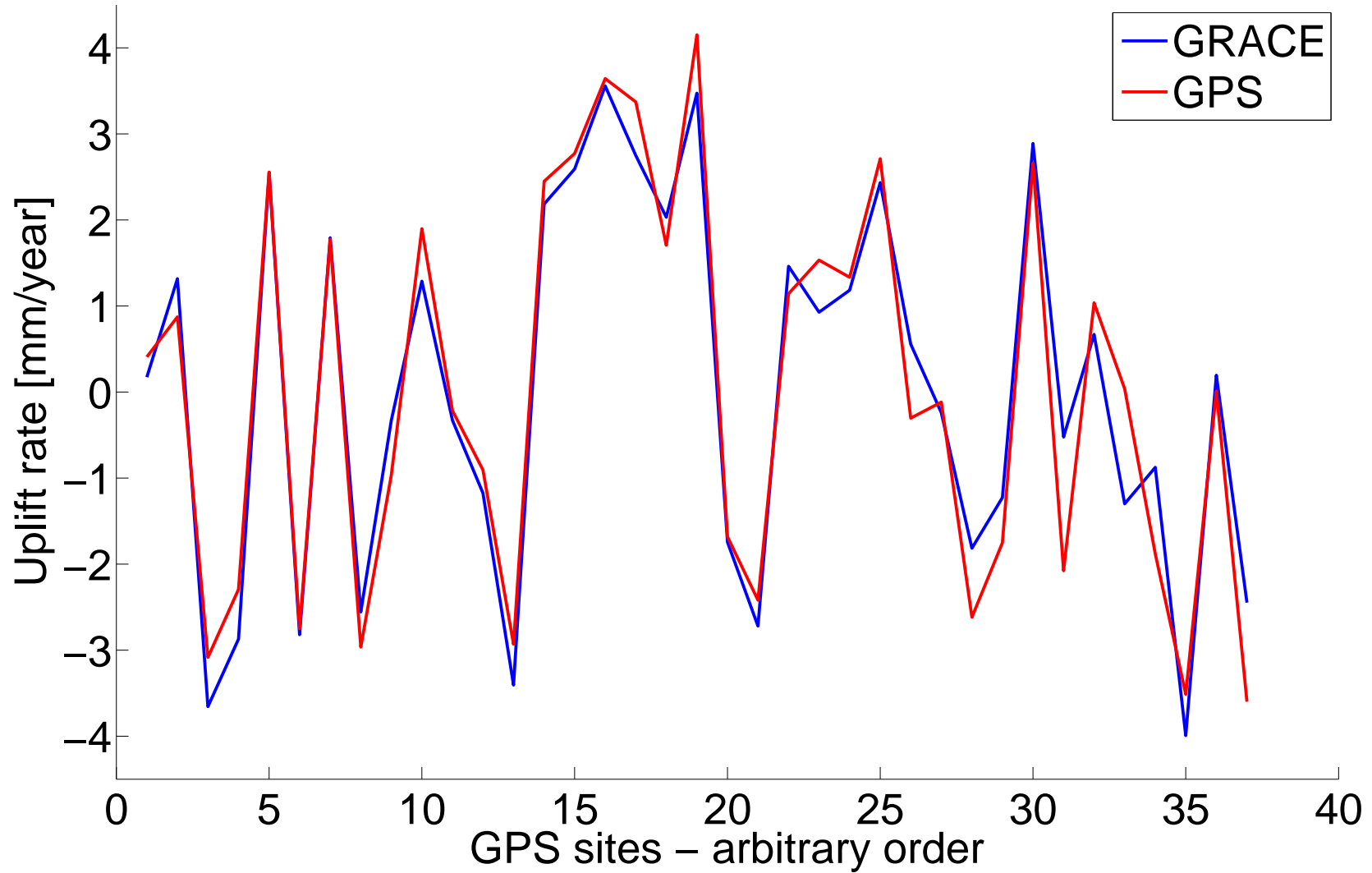


# Sensitivity kernels



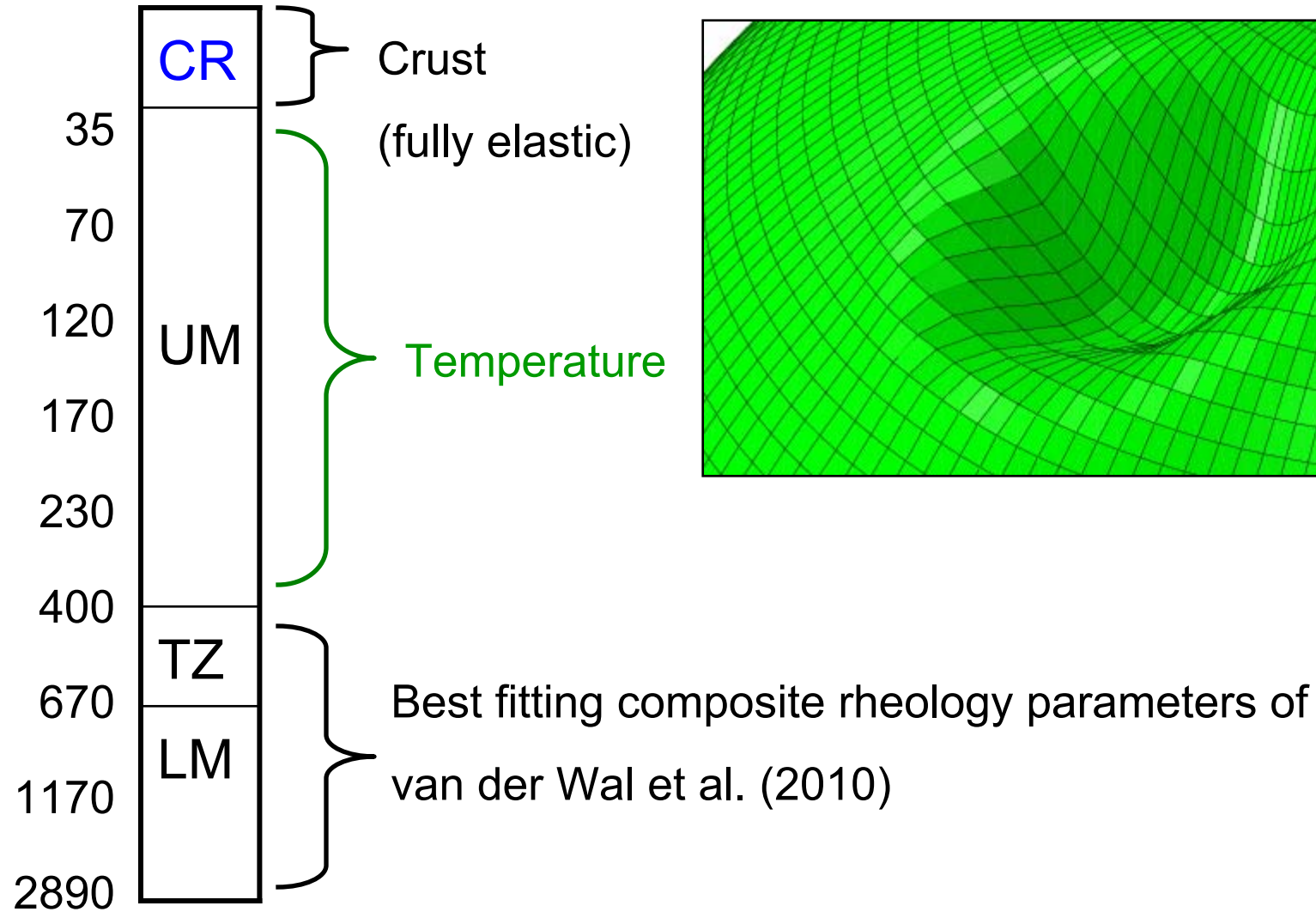
Gravity rate (uplift rate) in North America is more sensitive to the lower mantle viscosity

## After scaling





# Model



# Model

*Hirth & Kohlstedt (2003)*

$$\dot{\epsilon} = A_D \sigma^n d^{-p} fH_2O^r \exp(\alpha\phi) \exp\left(-\frac{E + pV}{RT}\right)$$

$A_D$  pre-exponent factor

$n$  stress exponent (3.5)

$d$  grain size (0.5–4 mm)  
**Kukkonen&Peltonen (1999)**

$fH_2O$  water fugacity

$\phi$  melt factor (0)

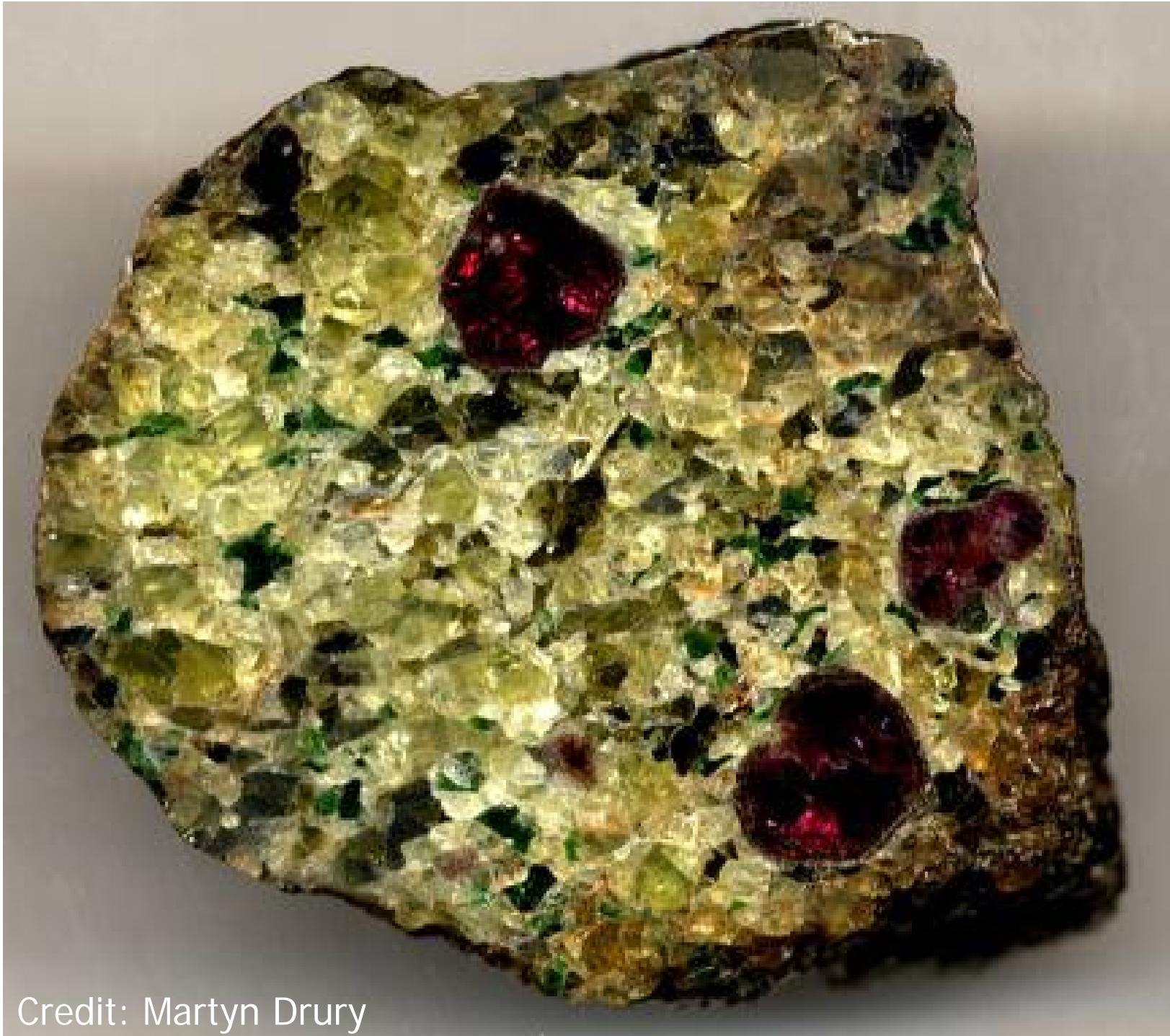
$E$  activation energy

$P$  pressure

$V$  activation volume

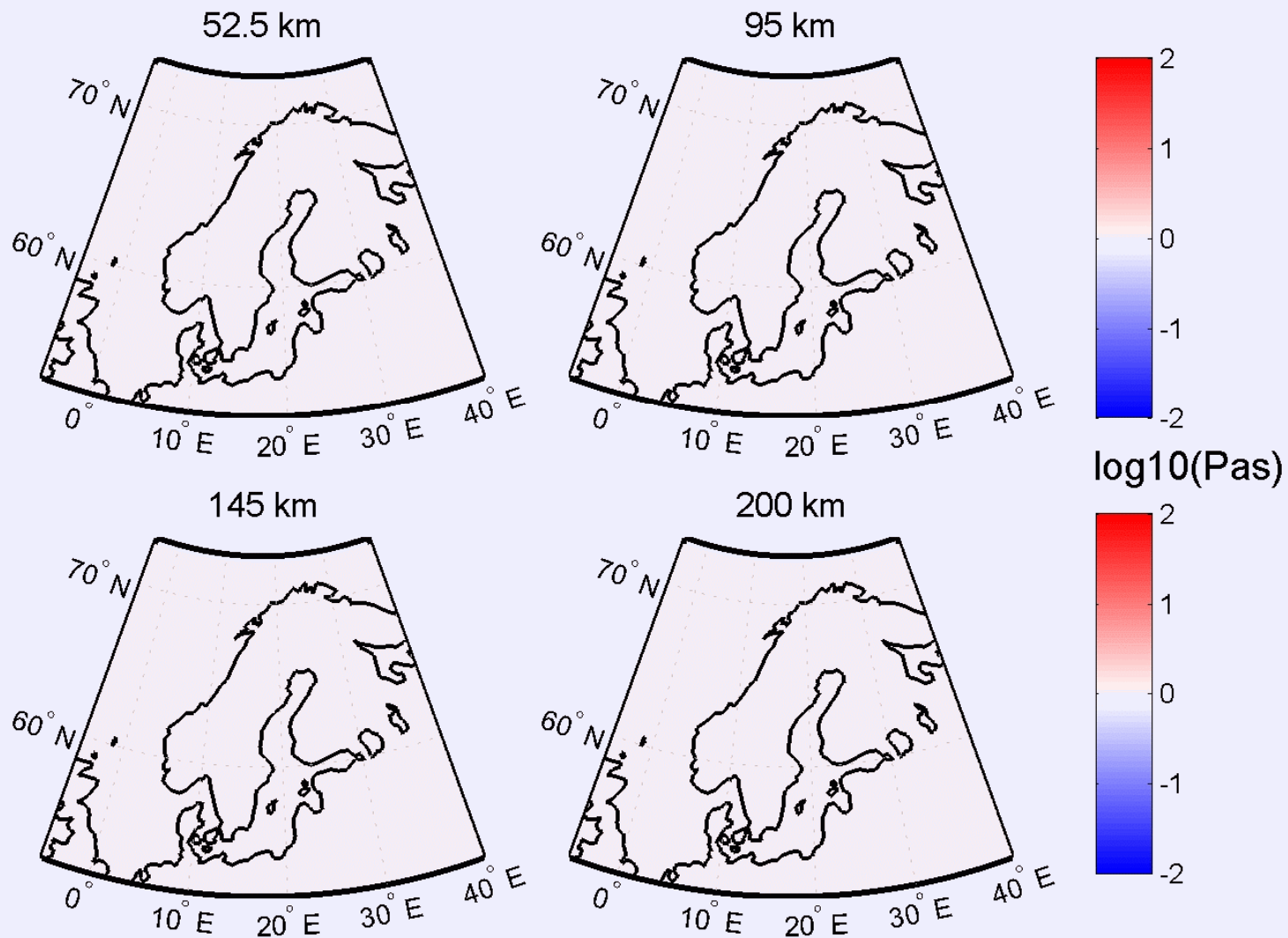
$T$  temperature

$R$  gas constant



Credit: Martyn Drury

30 ka B.P.



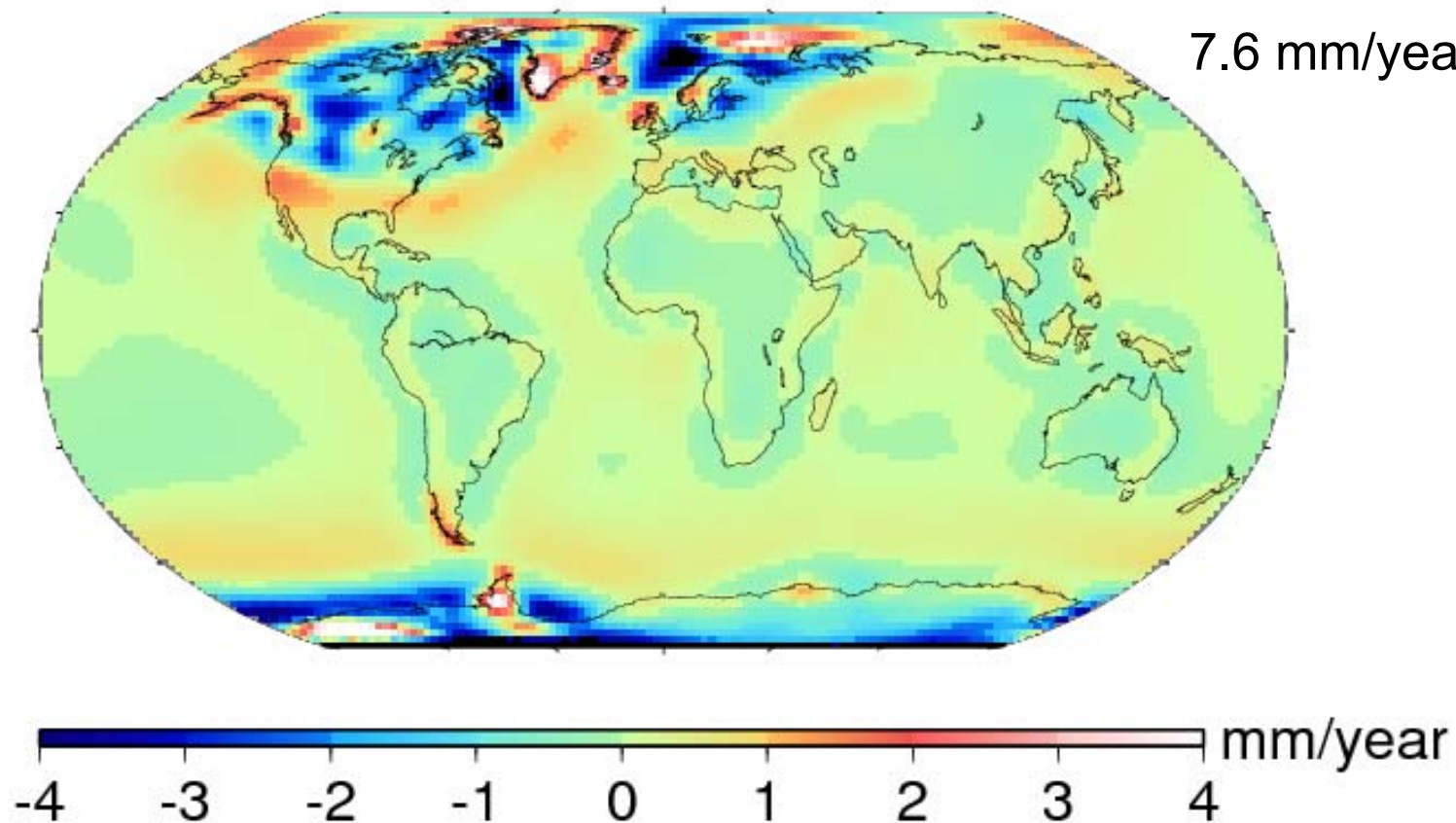
Barnhoorn, van der Wal, Drury, Vermeersen (G-cubed 2011)

# Uncertainty: 3D temperature

Temperature II – Temperature I

Max.

7.6 mm/year



van der Wal et al., (in prep.)